

Ratios of Extinction, Mass, Surface Area to Backscatter, and Mass and Surface area to Extinction derived from 30 years of mid latitude OPC measurements

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Motivation

The basic equations

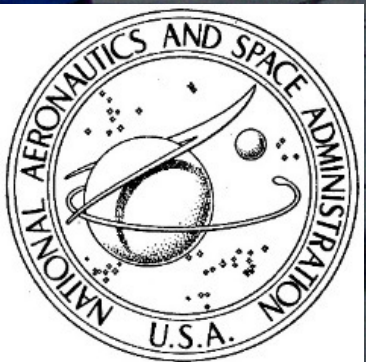
Profiles of temperature, weight fraction, index of refraction

Profiles of Backscatter, Extinction, Mass, Surface Area

Profiles of the ratios: E:B, M:B, SFA:B, M:E

Scatter plots to derive dependencies of E:B, M:B, SFA:B, M:E

Application to Tskuba Lidar data, from Osamu Uchino at NIES in Japan



The OPC data presented here are publicly available at <https://doi.org/10.15786/21534894>



Motivation

- E/B ratio - the Lidar ratio
 - Needed for lidar measurement retrievals
 - Can be used to convert lidar profiles to extinction profiles, and aerosol optical depth, e. g. Ridley et al., 2014, GRL, 41, 7763–7769
- m/B, m/E ratios
 - Allows optical measurements to be used to estimate aerosol mass, useful for estimating aerosol lifetimes in the stratosphere
 - With assumptions integrated profile measurements (a typical lidar data product) can be used to estimate the stratospheric sulfur burden
- s/B, s/E ratios
 - Allows optical measurements to be used to estimate aerosol surface area, important for stratospheric chemistry
 - The bad news these relationships are much less solid than the other ratios

We all know the basic equations for surface area, s , mass, m , extinction, E , and backscatter, B

$$s = \int_0^{\infty} \pi a^2 dn(a)/da da,$$

$$m = \int_0^{\infty} \pi a^3 /6 dn(a)/da da,$$

$$E = \int_0^{\infty} \pi a^2 Q_{ext}(\lambda, m, a) dn(a)/da da,$$

$$B = \left[\int_0^{\infty} \pi a^2 Q_{bks}(\lambda, m, a) \frac{dn(a)}{da} da \right] / 4\pi,$$

for aerosol size distribution $dn(a)/da$, where a is particle diameter, $Q_{ext/bks}$ the Mie extinction and backscatter cross sections, at wavelength, λ , and particle index of refraction, m .

Aside from the Mie calculations, the tricky part is getting index of refraction, which for sulfate is a function water vapor and temperature as these affect the weight fraction and solution density, both important for m .

But we know how to do this, e.g. Deshler et al., 2019, JGR, 124(9), 5058–5087

Extinction/backscatter calculations from in situ measurements

$$\beta_{\lambda}(m) = \int_0^{\infty} \pi a^2 \cdot Q(a, \lambda, m) \cdot dn/d \ln(a) \cdot d \ln(a),$$

$Q(a, \lambda, m)$ Is the extinction cross section, with m the index of refraction, the only tricky part of the calculation.

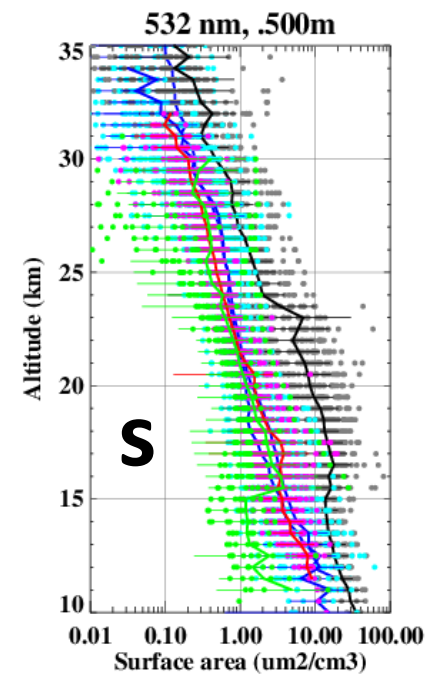
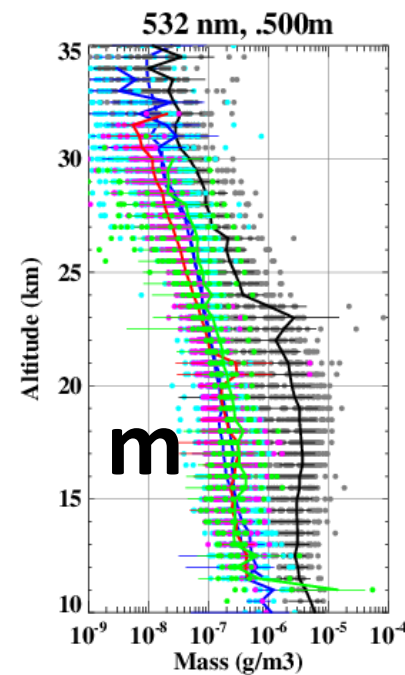
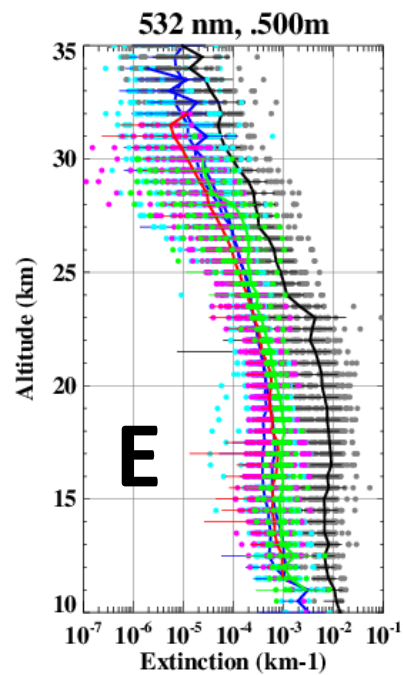
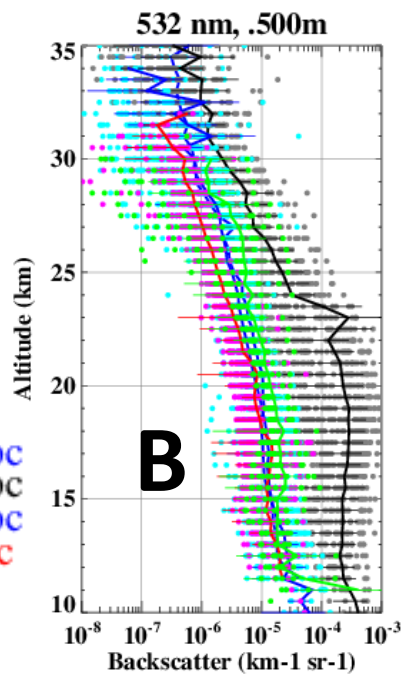
- Calculation of index of refraction, m , requires:
 - Water vapor (assumed) and temperature (measured) profiles
 - Sulfuric acid weight percent = $f(\text{wv}, T)$ [Steele and Hamill, 1981]
 - Palmer and Williams [1975] for $m(300 \text{ K}) = f(\text{sulfuric acid weight percent})$
 - Lorentz-Lorenz for index of refraction = $f(T, \text{density of sulfuric acid}, \rho(T))$

$$m(T) = \sqrt{\left(\frac{1 + 2A\rho(T)}{1 - A\rho(T)} \right)}, \text{ where } A = \frac{m^2(300) - 1}{[m^2(300) + 2]\rho(300)}$$

$\rho(T)$, is calculated from Luo et al. (1996),

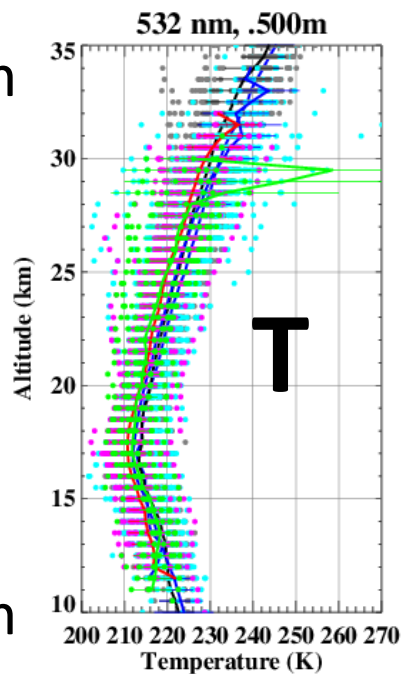
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163 OPC profiles

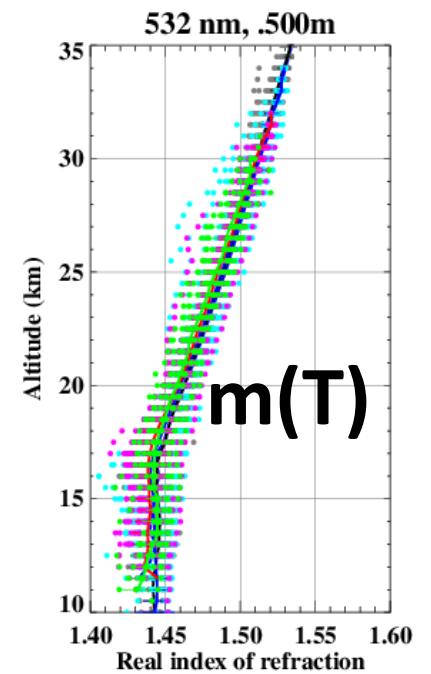
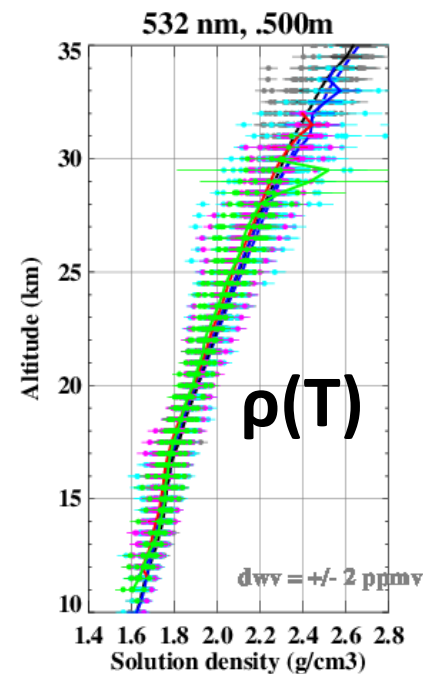
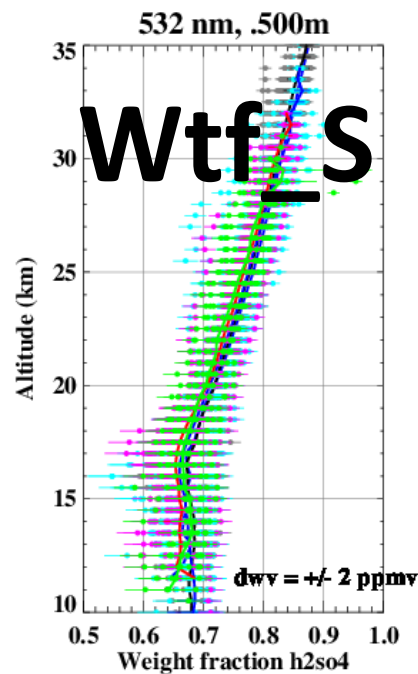


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30 km

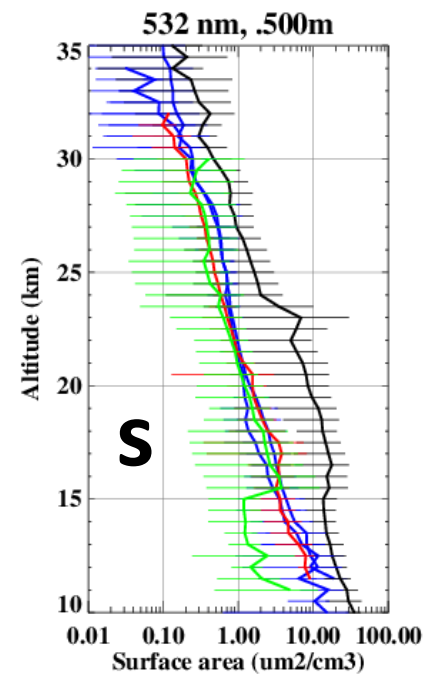
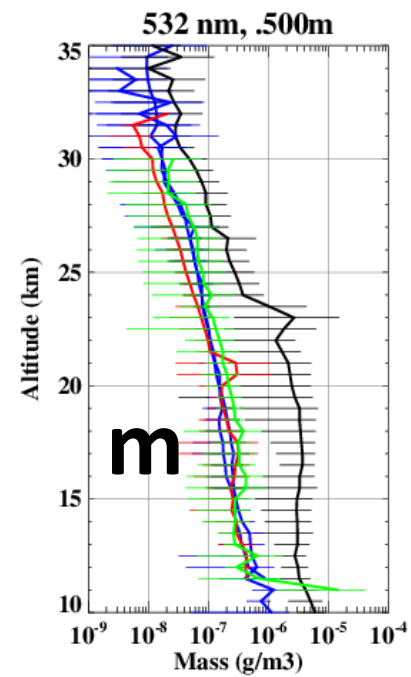
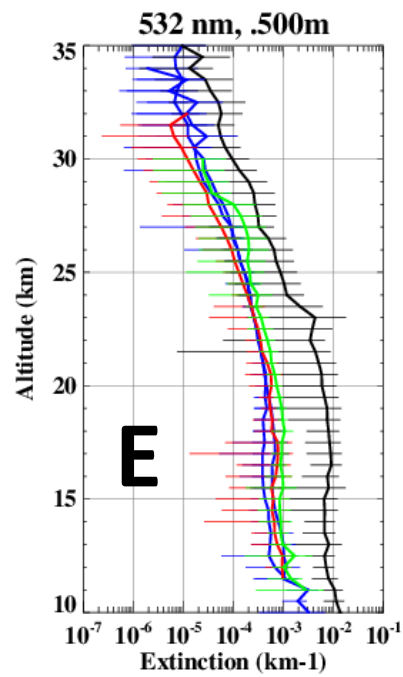
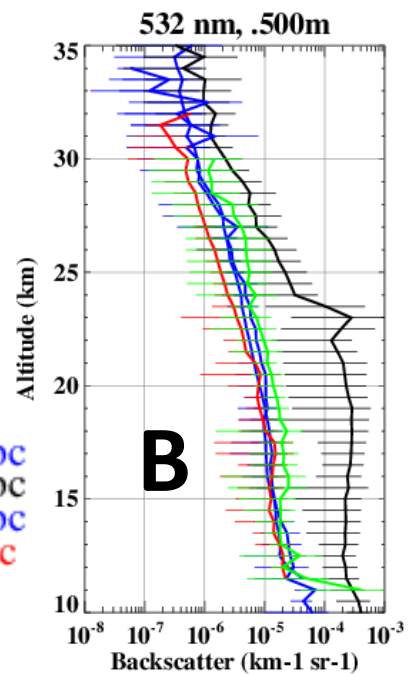


10 km



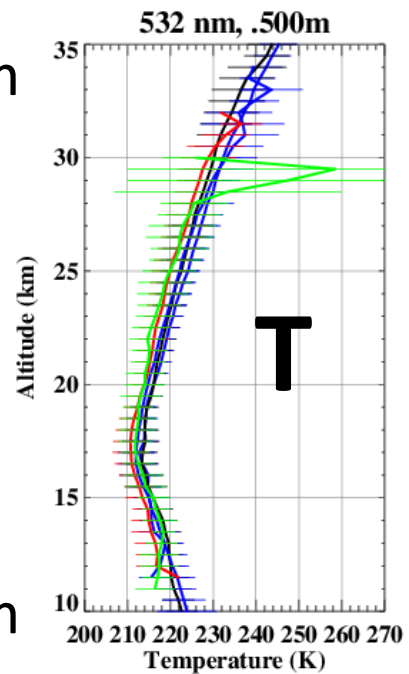
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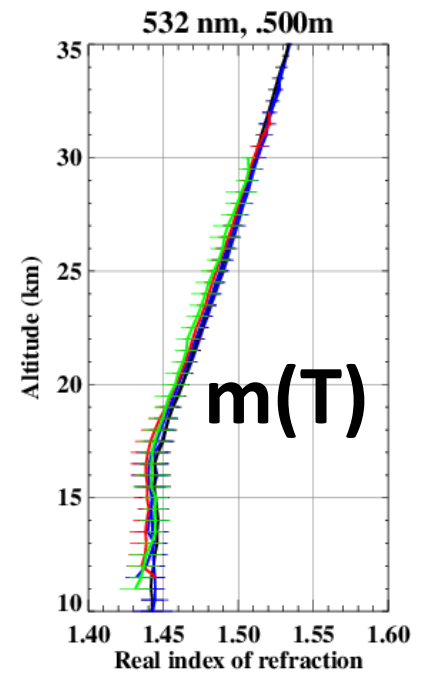
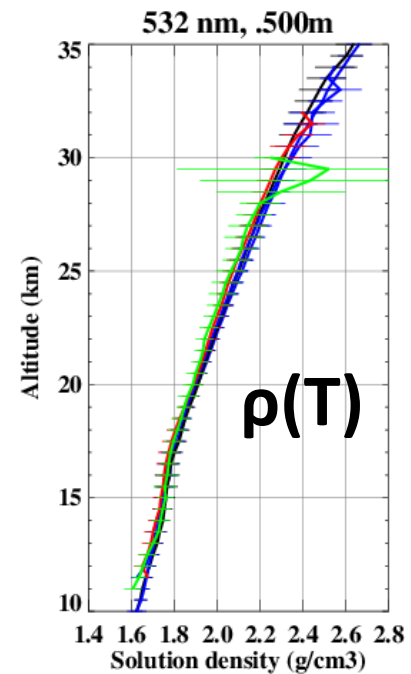
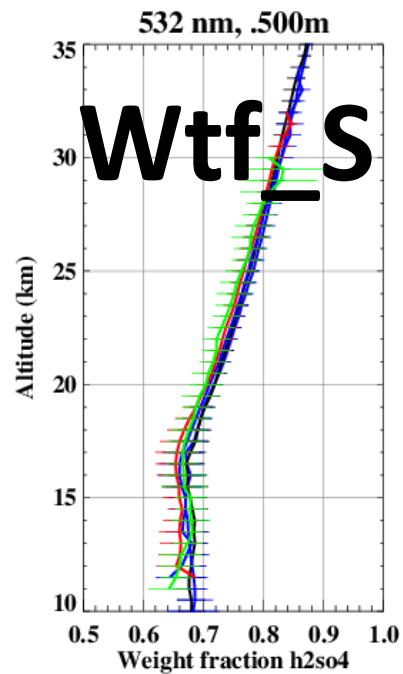


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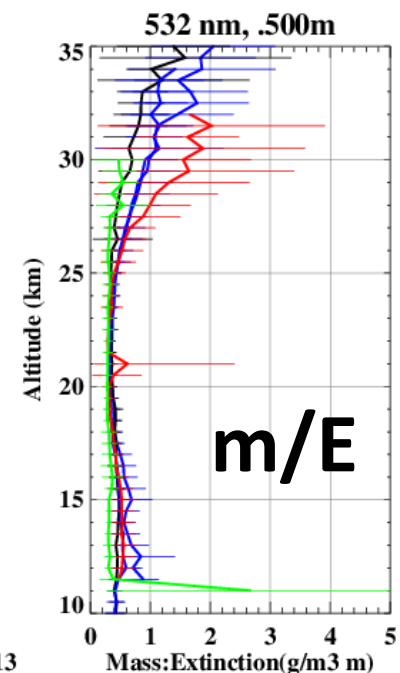
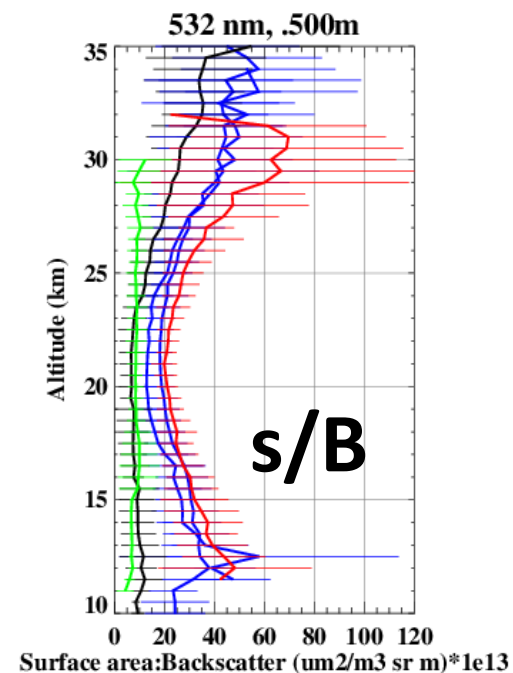
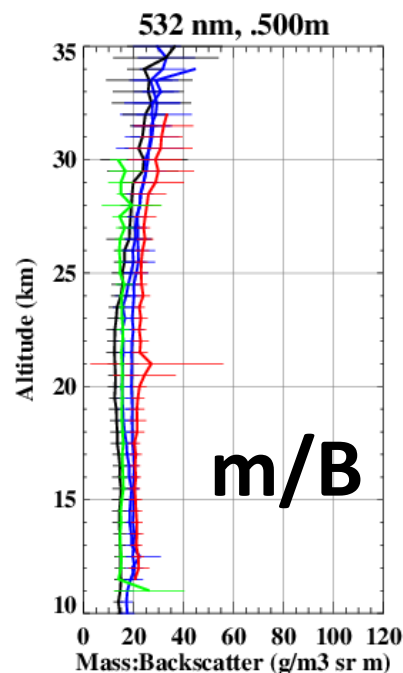
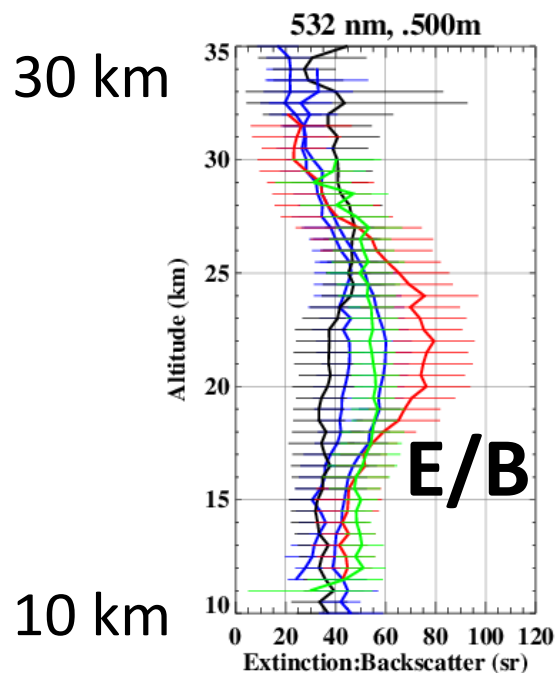
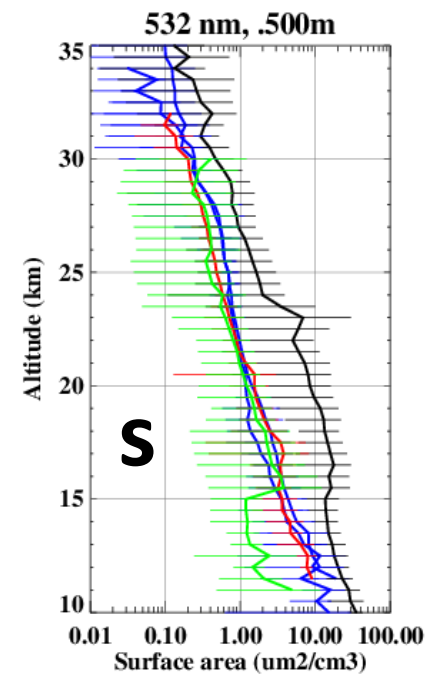
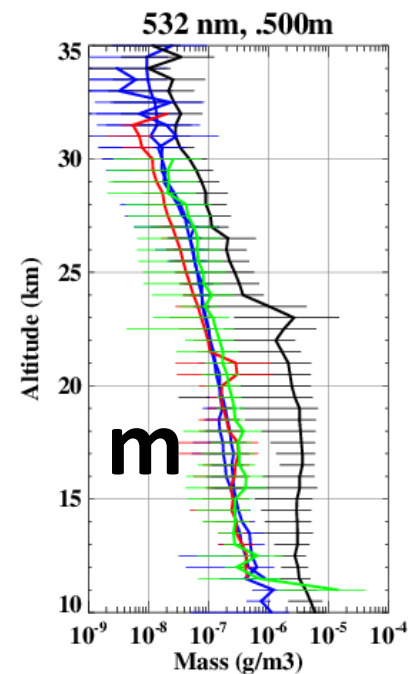
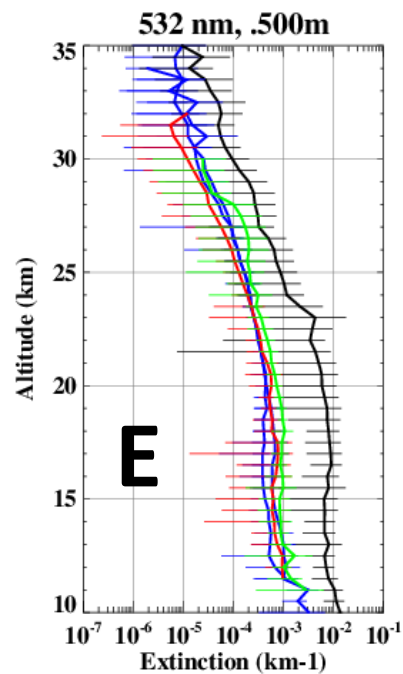
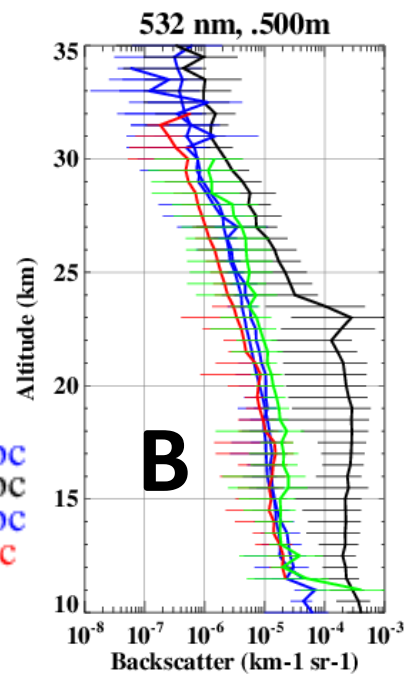
10 km



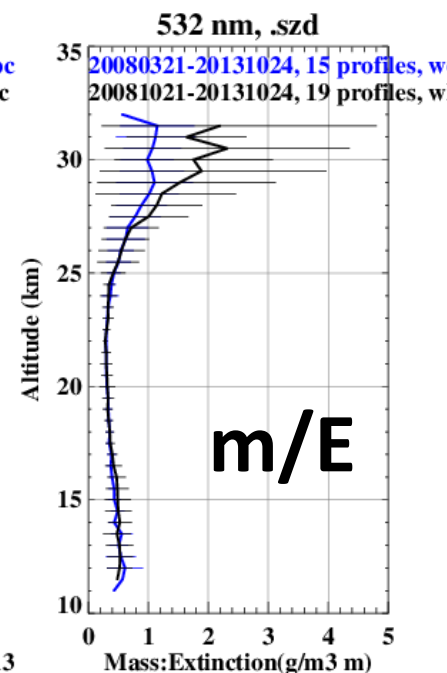
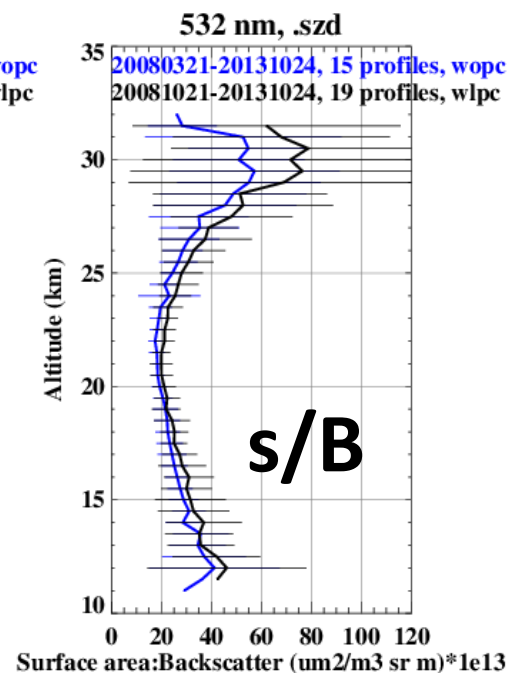
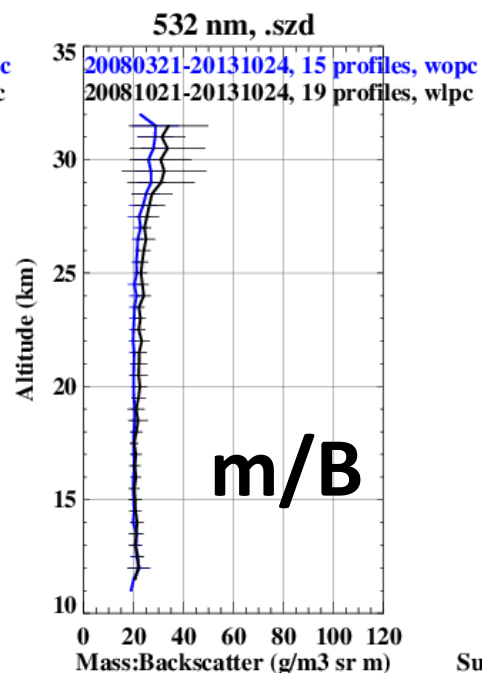
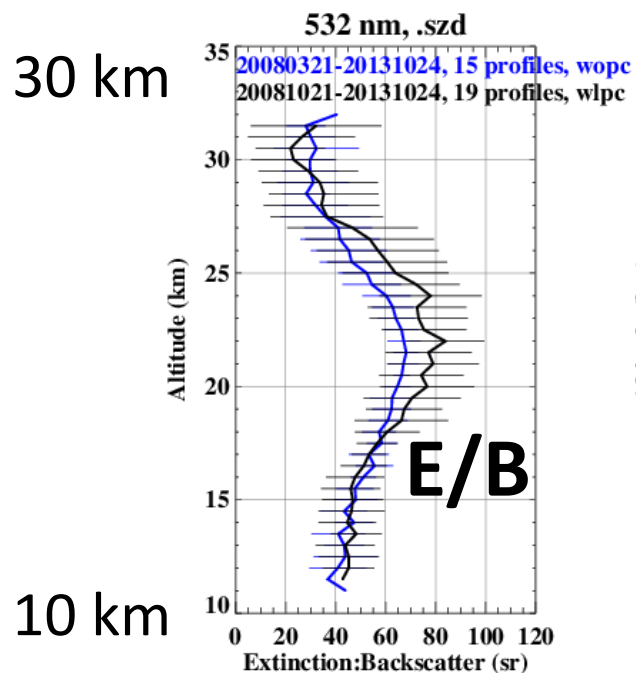
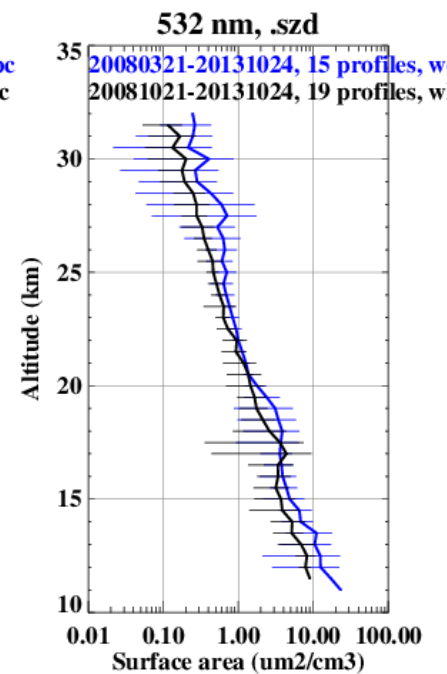
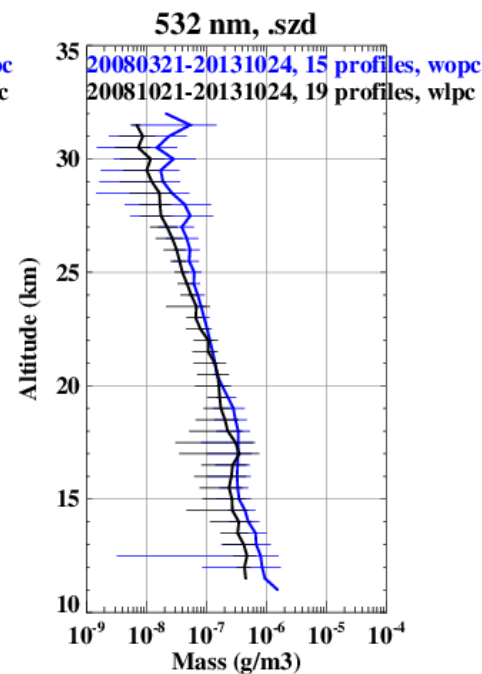
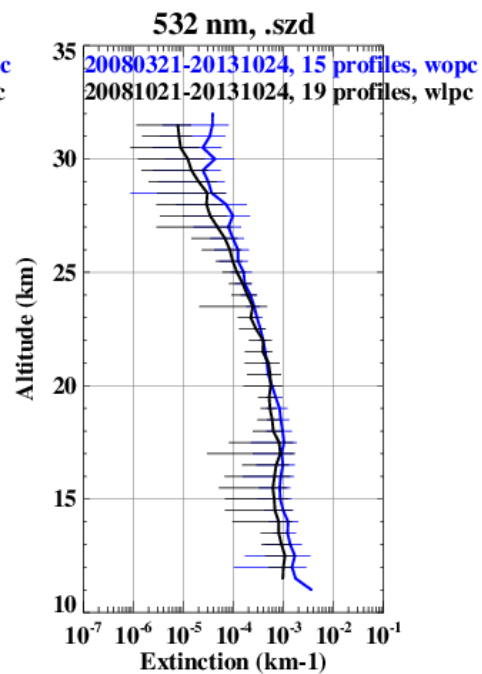
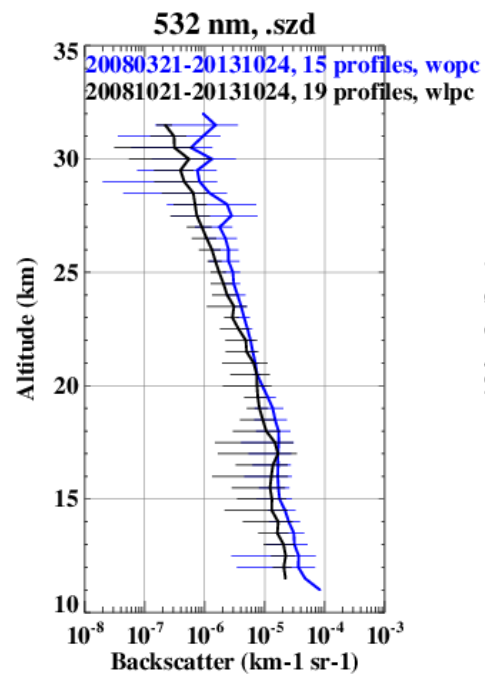
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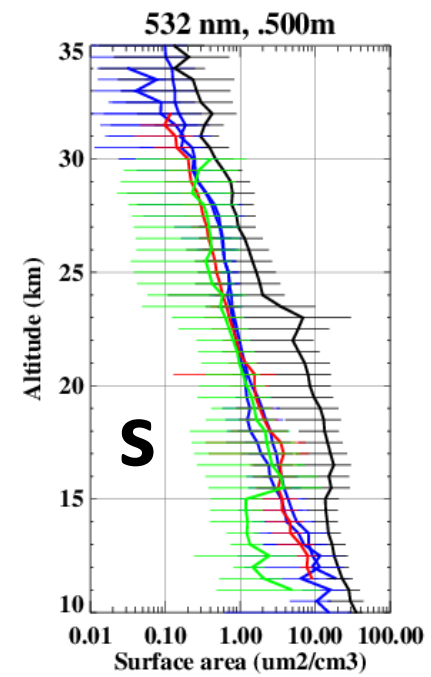
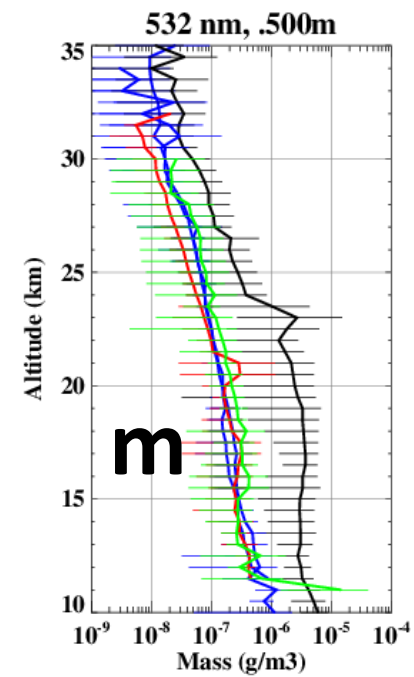
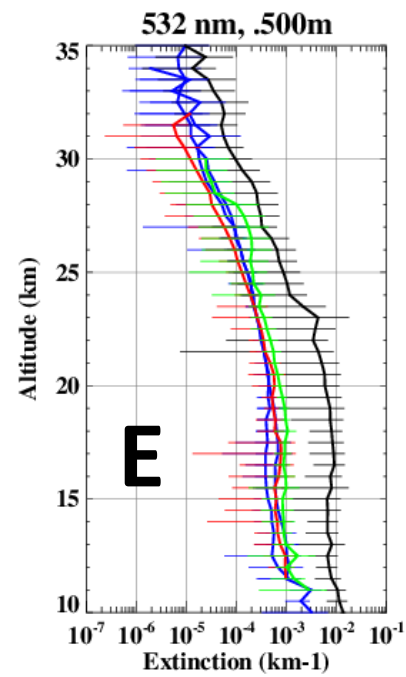
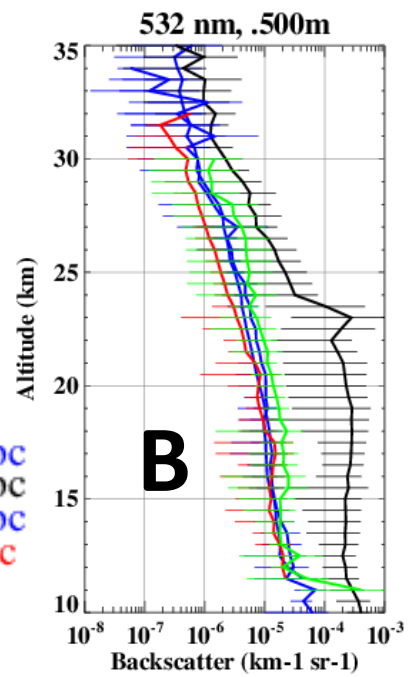


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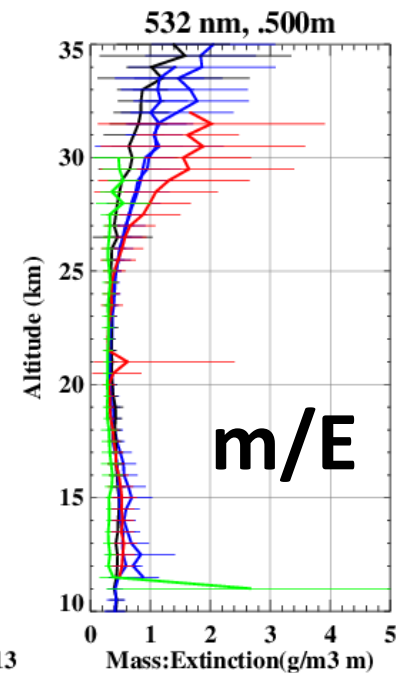
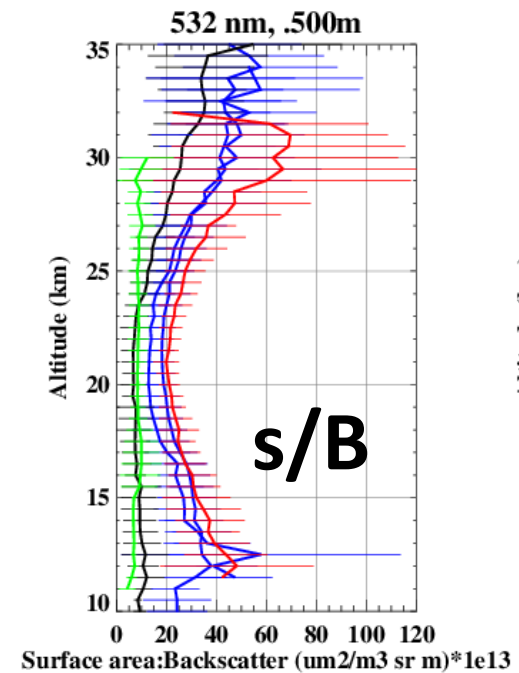
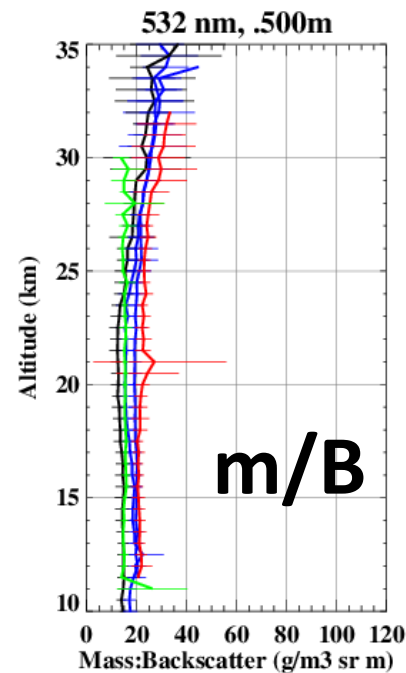
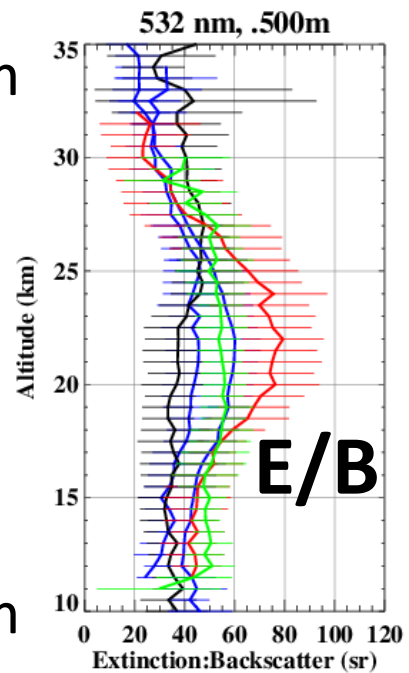


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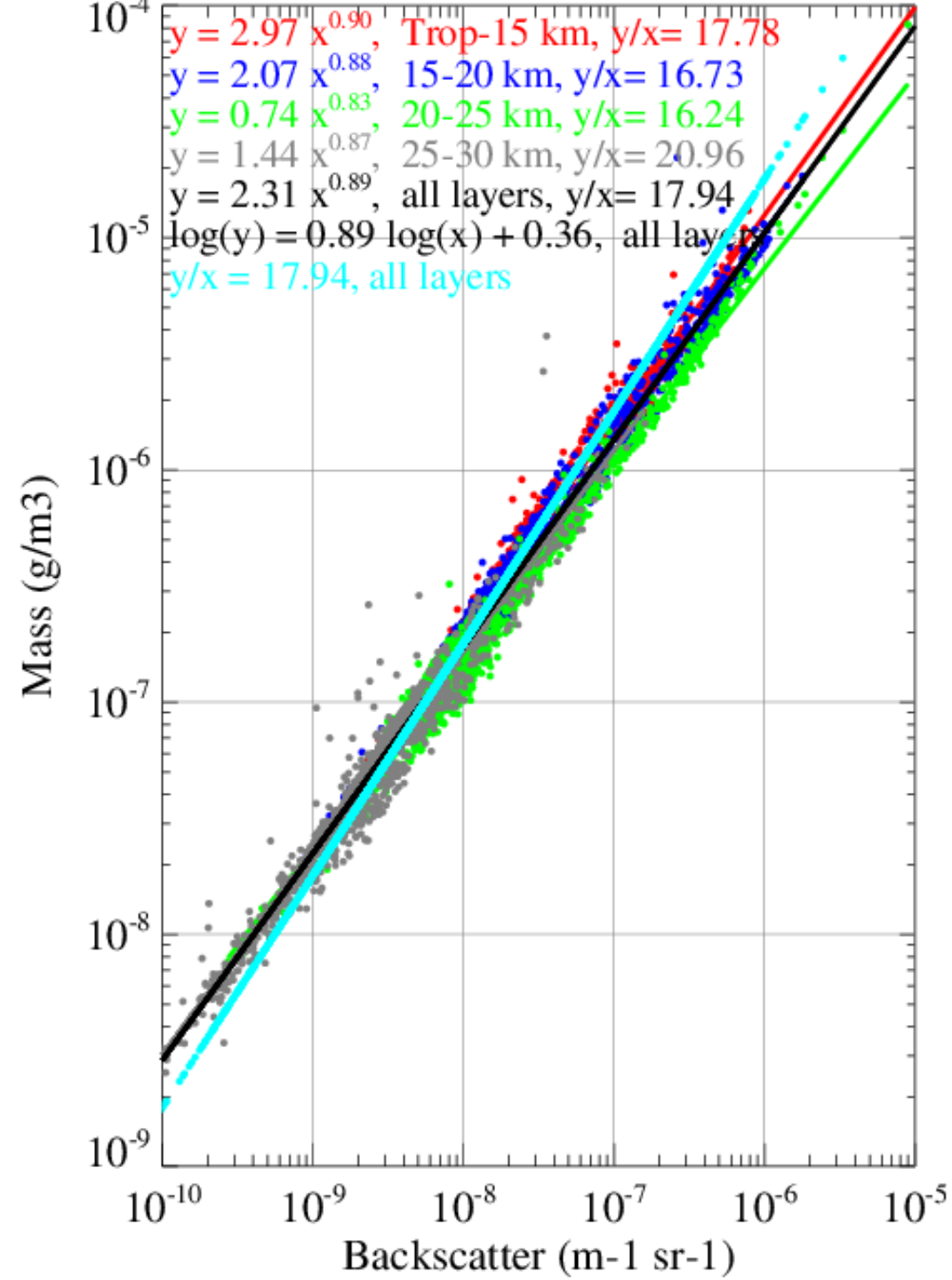
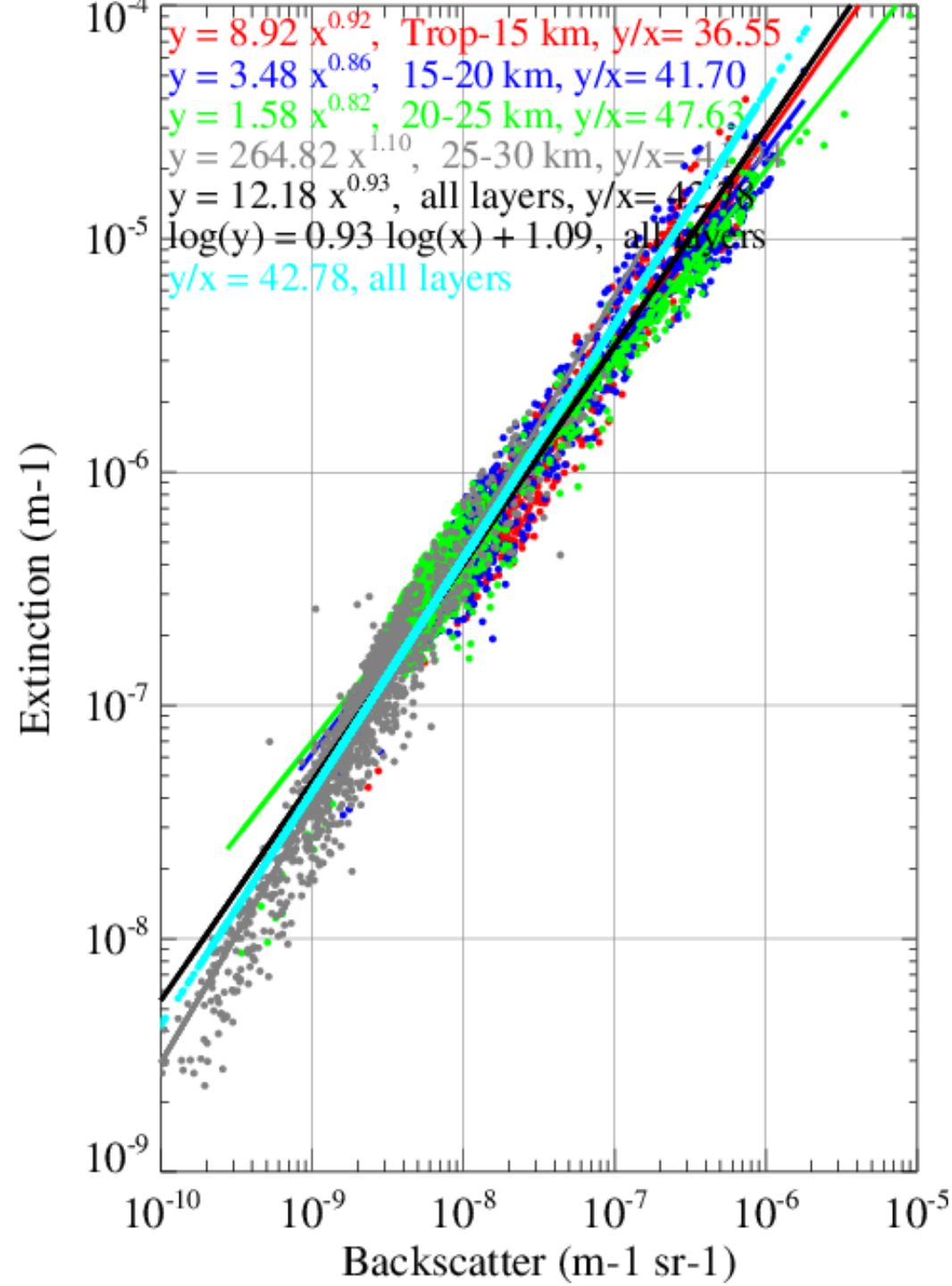
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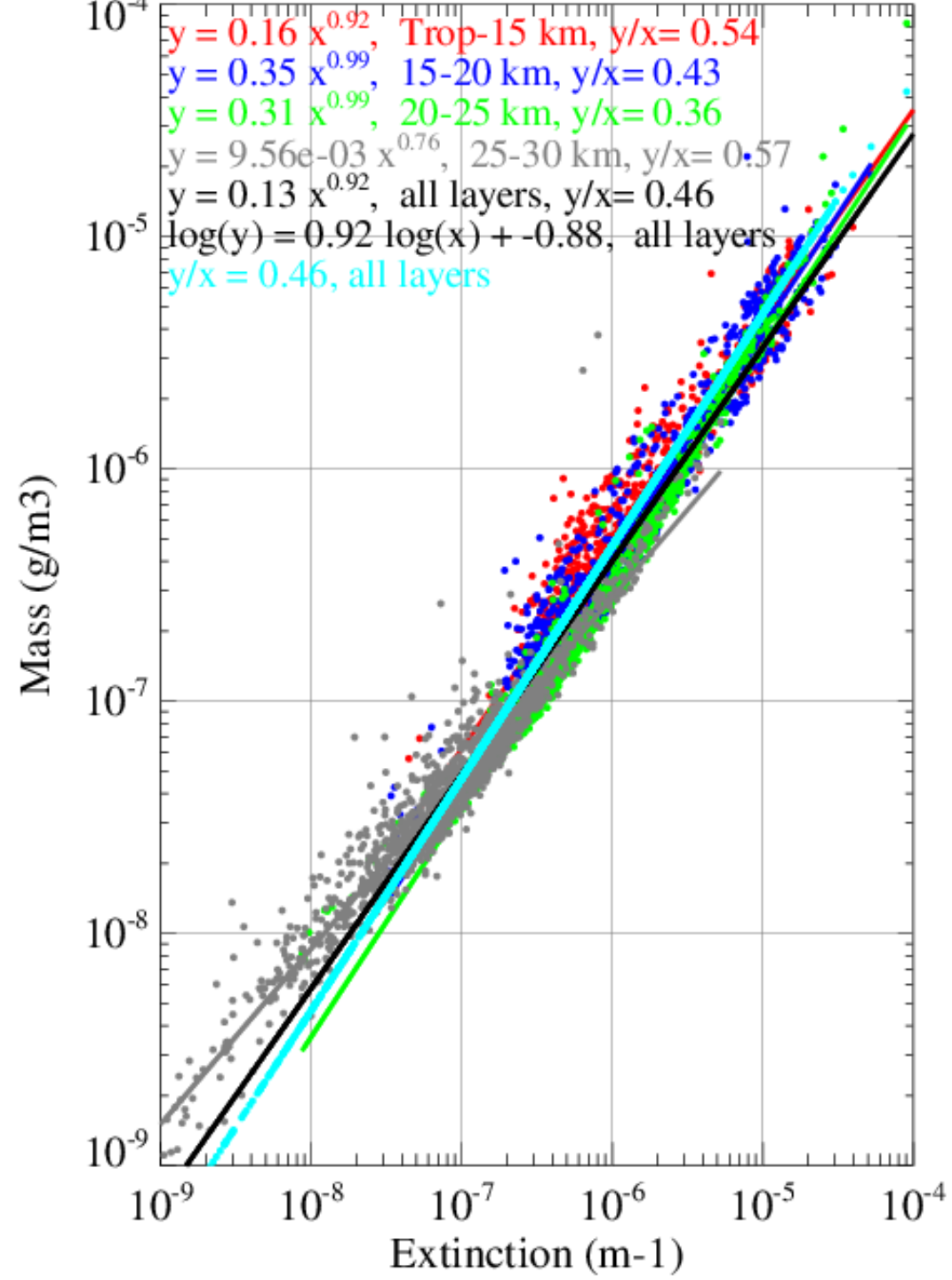
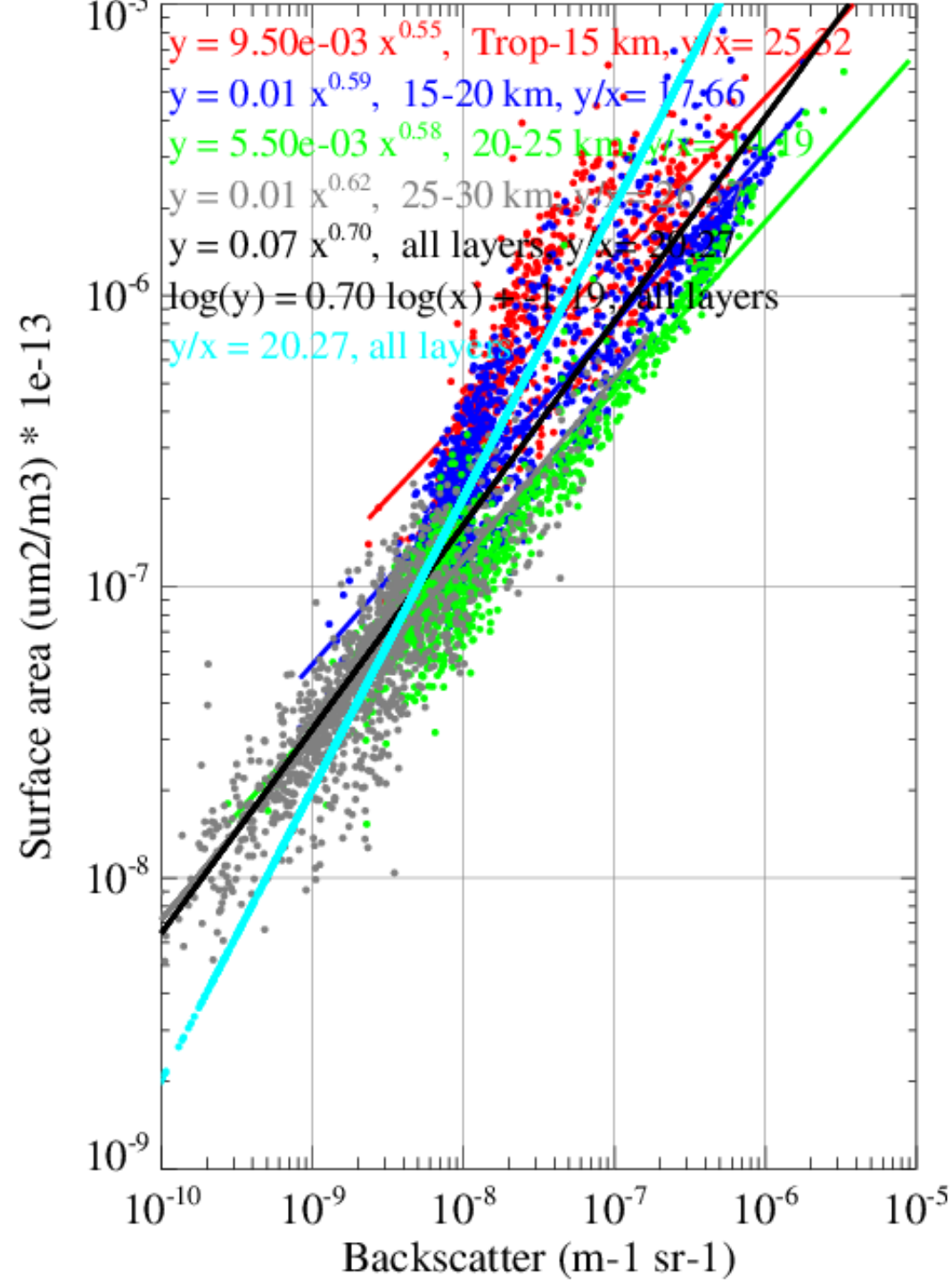


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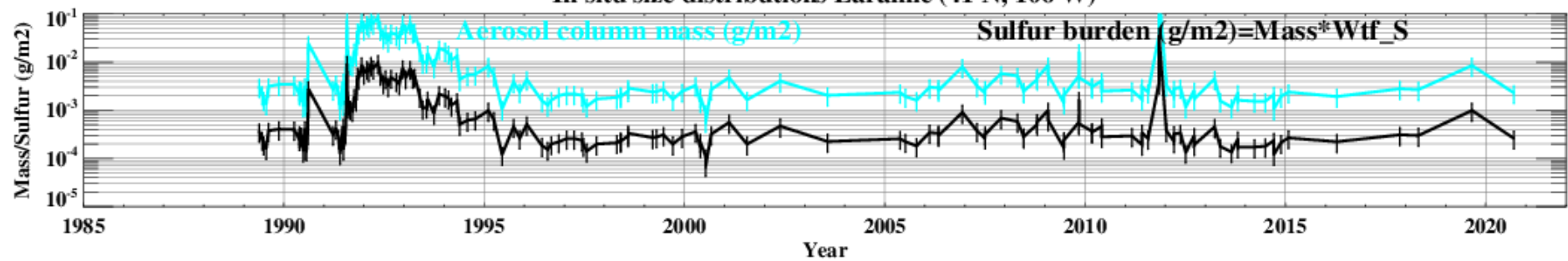


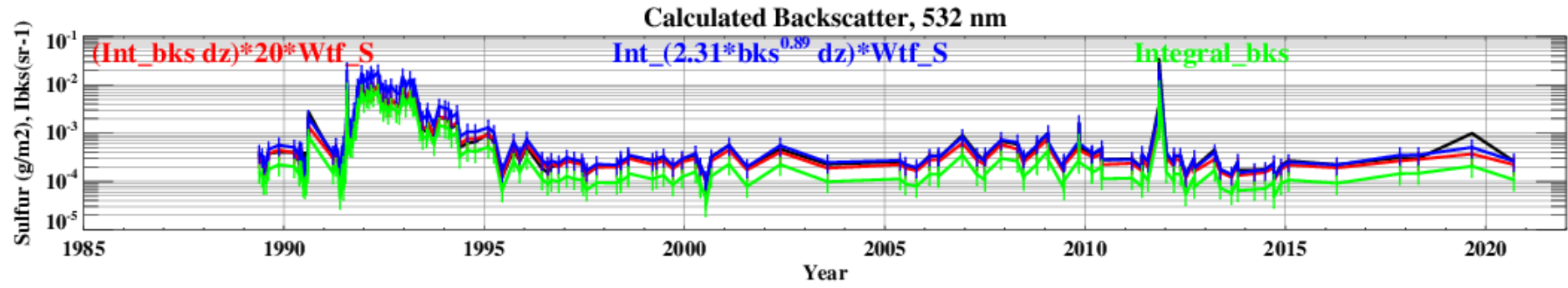
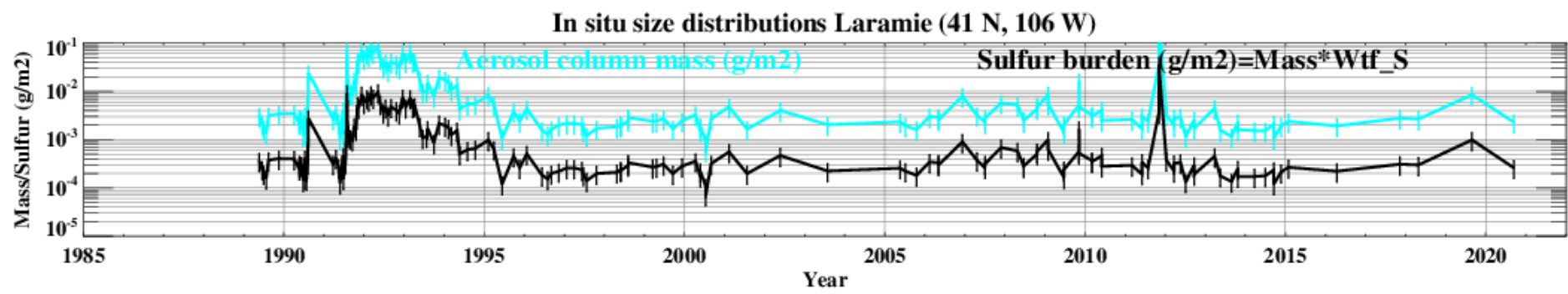
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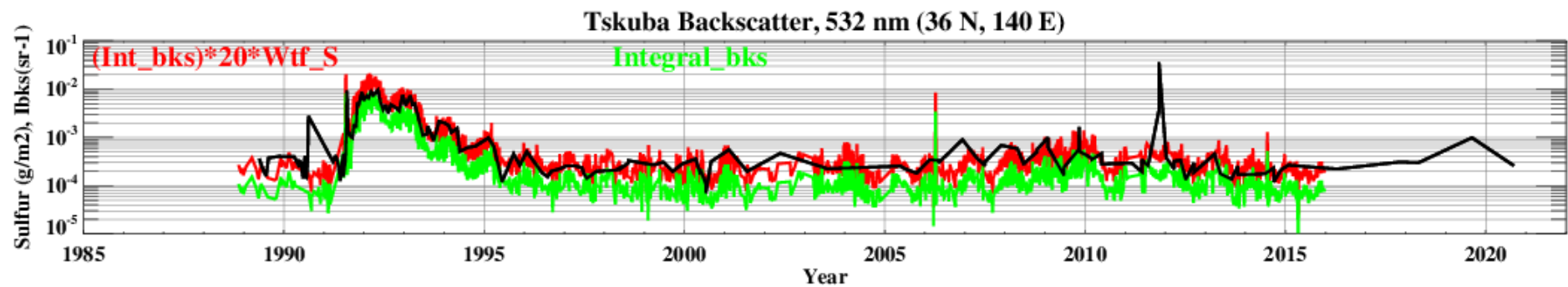
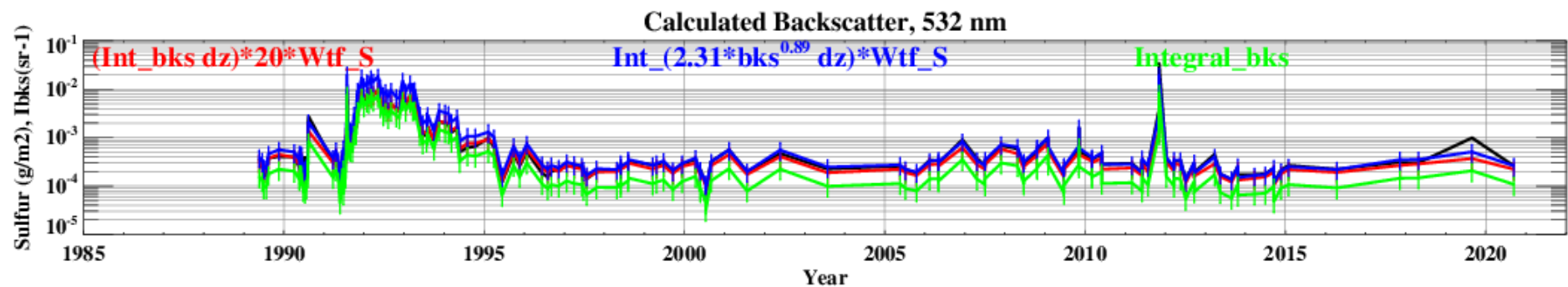
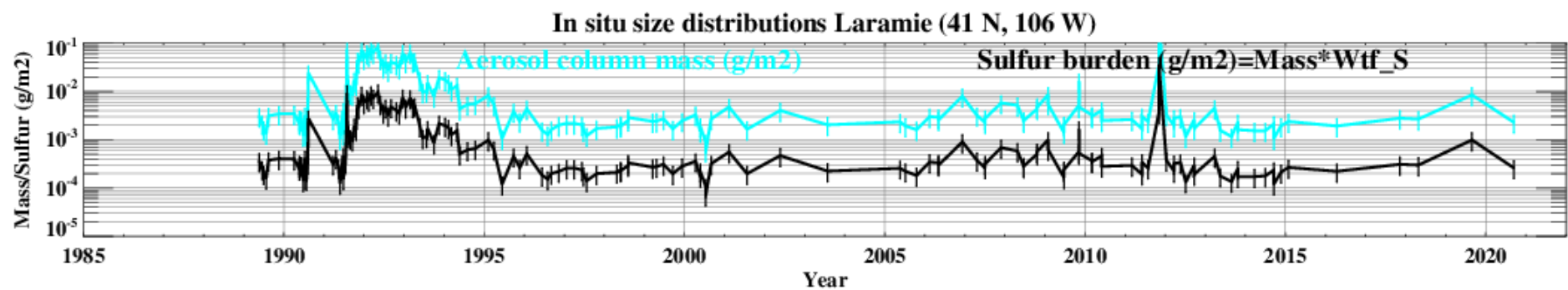




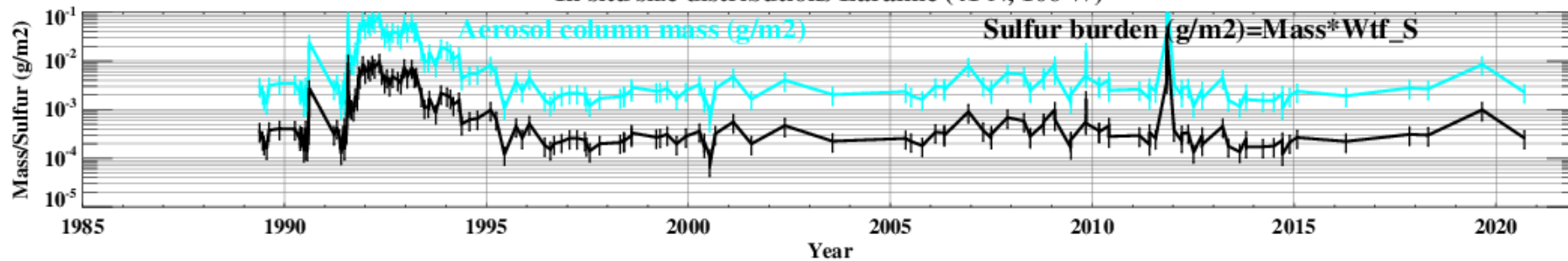
In situ size distributions Laramie (41 N, 106 W)



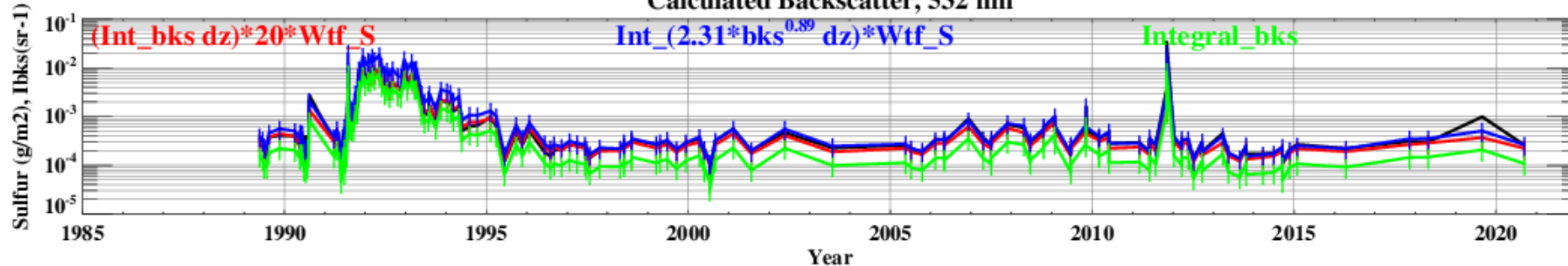




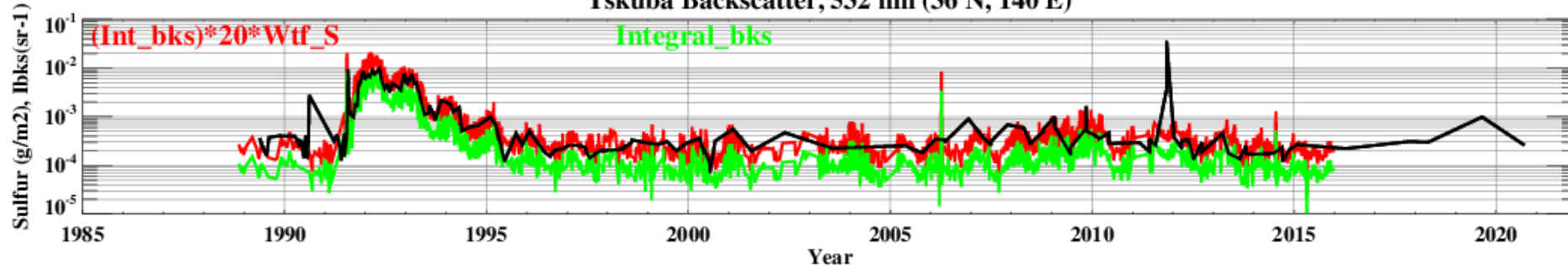
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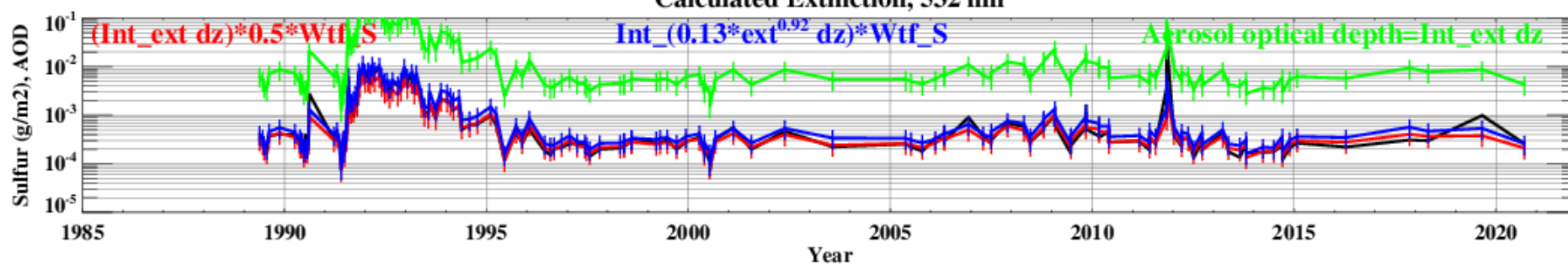
Calculated Backscatter, 532 nm

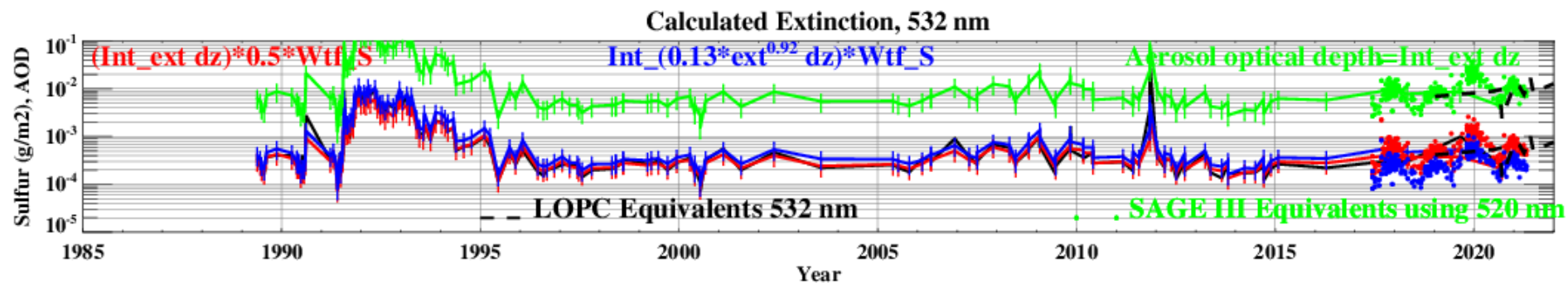
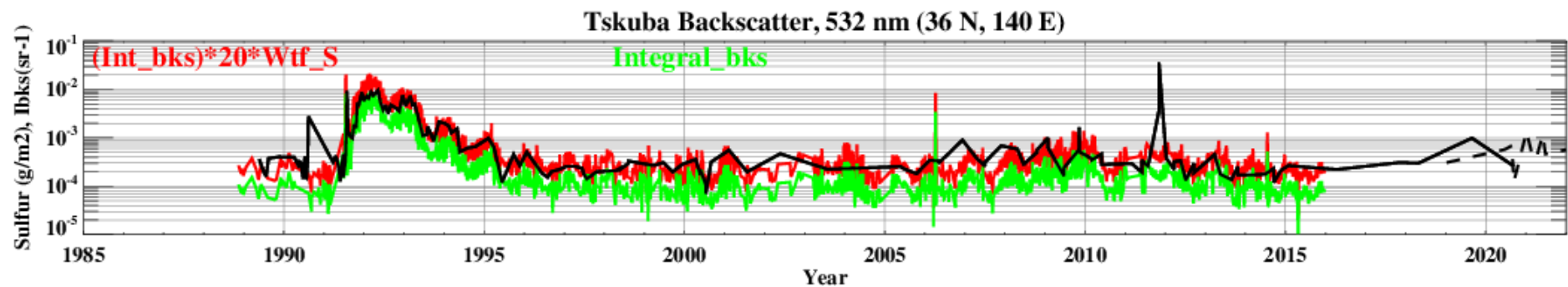
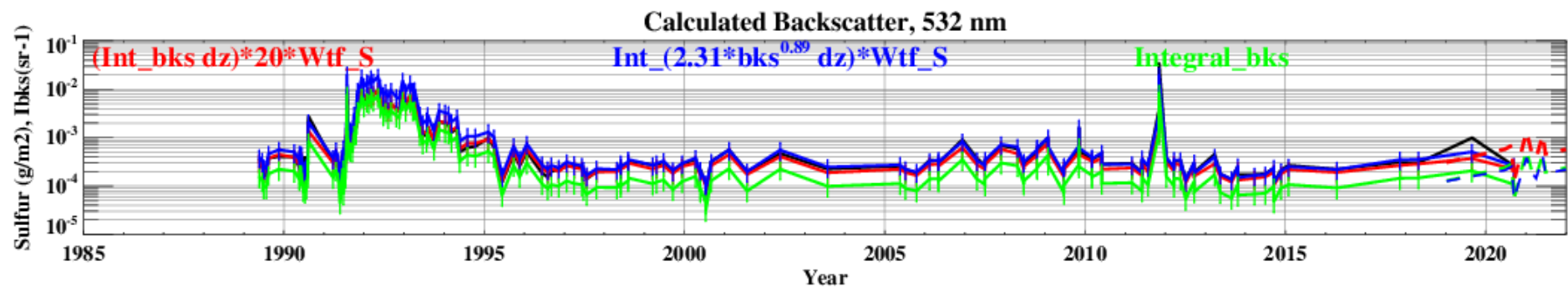
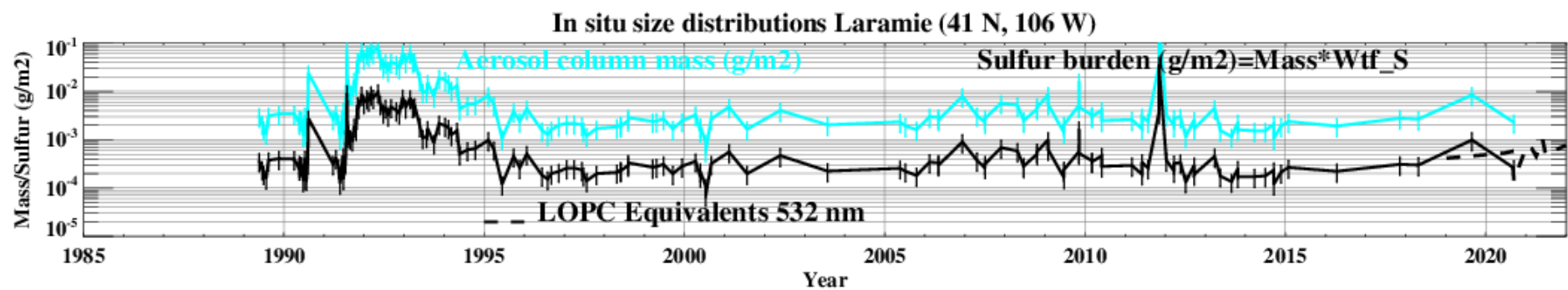


Tskuba Backscatter, 532 nm (36 N, 140 E)



Calculated Extinction, 532 nm





Conclusions

- Distribution moments (backscatter, extinction, mass, surface area) from three generations of OPCs are consistent, outside of large geophysical events, across 30 years of measurements.
- Ratios of these moments provide extinction or mass from backscatter measurements, or mass from extinction measurements.
- These relationships are reasonably tight in spite of altitude and temporal dependencies, and can often be approximated with a simple ratio.
- Relationships to surface area are much more variable. This is likely due to the dependence of surface area on smaller particles which don't contribute significantly to the optical quantities.
- Application of the mass:backscatter ratio to Tskuba lidar data showed surprisingly good correspondence of aerosol sulfur burden estimates between the Wyoming (OPC) and Japanese (Lidar) measurements, both in magnitude and temporally.
- These OPC data are publicly available at <https://doi.org/10.15786/21534894>